

Modelling non-stationary spatially-dependent extremes

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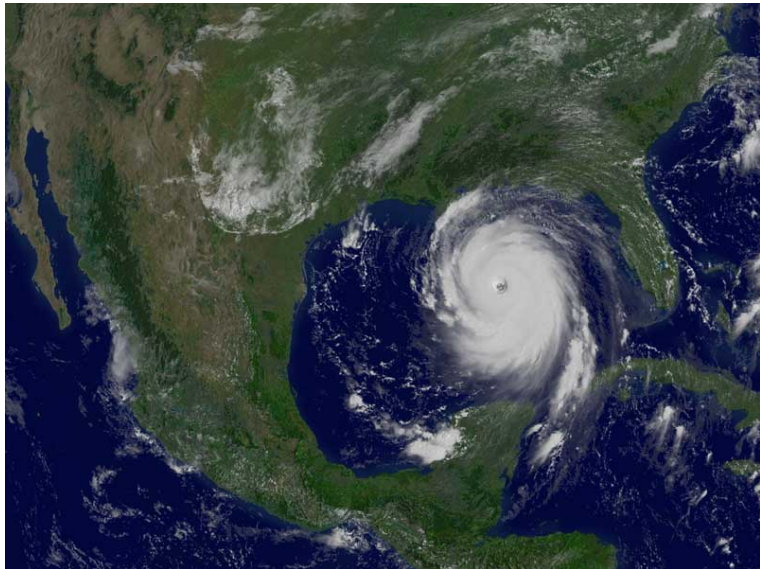
Newcastle University
18th November 2011

This is joint work with Philip Jonathan (Shell Research)

- Wave height data → design of safe marine structures.
- Extreme value modelling
- Modelling issues
 - Spatial non-stationarity and dependence
 - Thresholds for non-stationary extremes
 - Model parameterisation
- Theoretical and simulation studies
- Wave height data

- Hindcasts of Y **storm peak significant wave height** (in metres) in the Gulf of Mexico.
 - **wave height** : trough to the crest of the wave.
 - **significant wave height** : the average of the largest 1/3 wave heights. A measure of sea surface roughness.
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 - **storm peak**: largest value from each (hurricane-induced) storm.
- a 6×12 grid of **72 sites** (≈ 14 km apart).
- Sep 1900 to Sep 2005 : **315 storms** .
- average of 3 observations (storms) per year, at each site.







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- Extrapolation !

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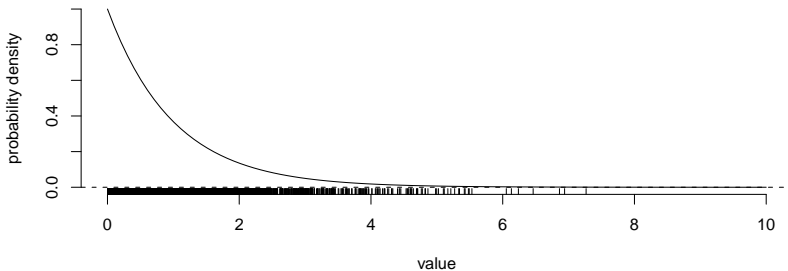
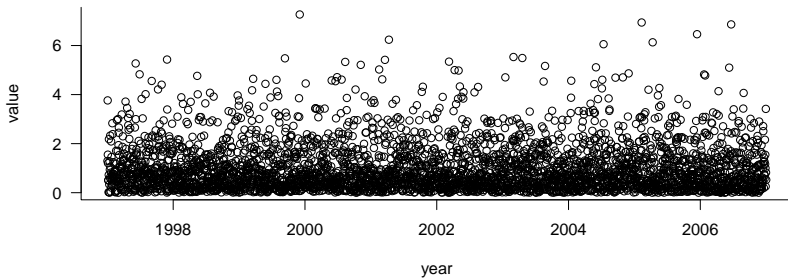
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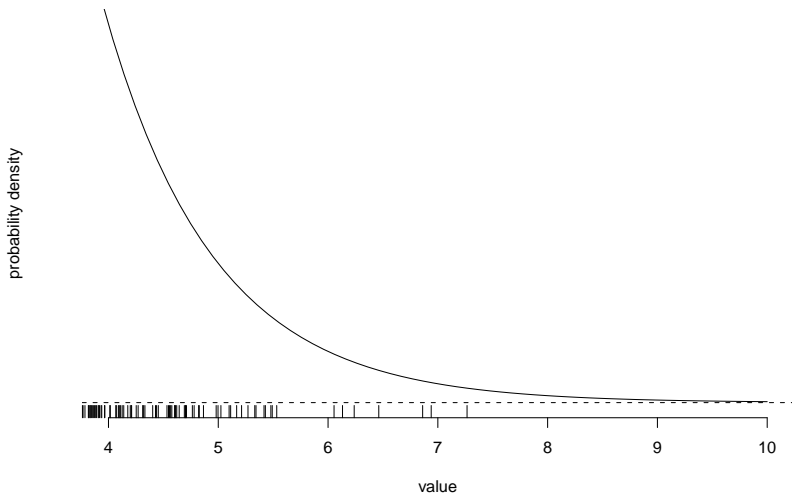
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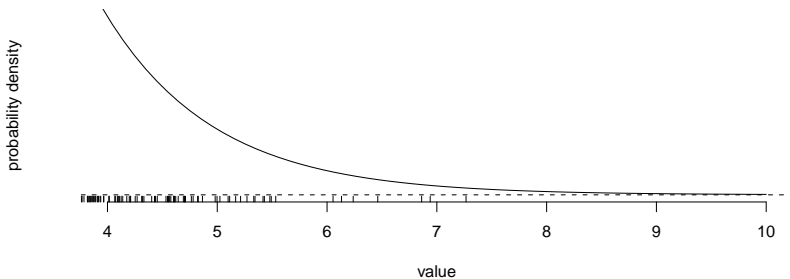
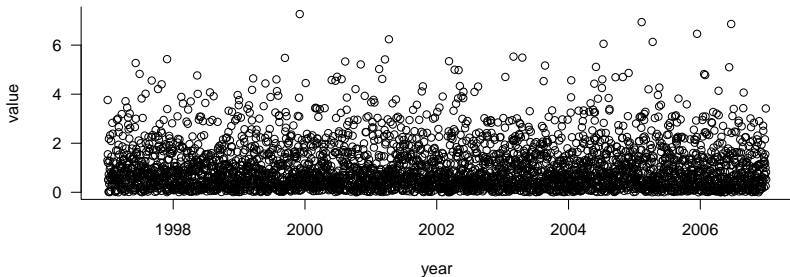
- Base inferences about future extreme behaviour on **extreme** data.

...but how should we define 'extreme data', and what models should we fit to these data?

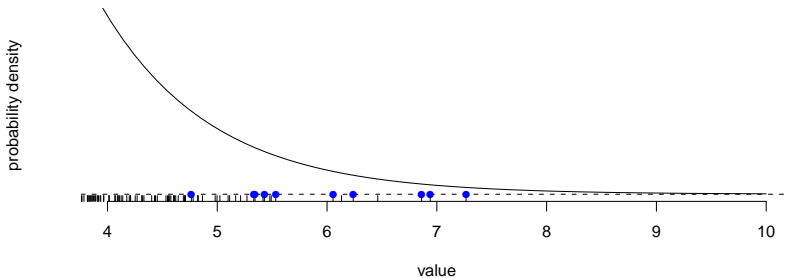
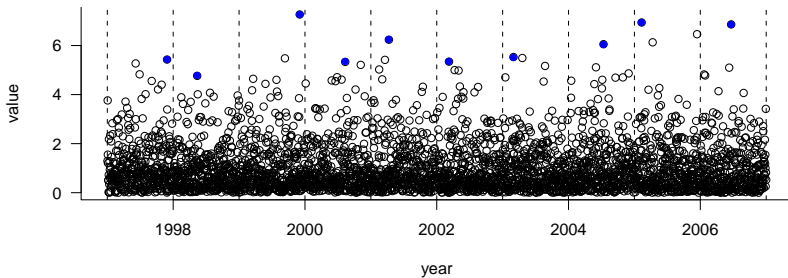
Simulated data [i.i.d. $\exp(1)$]



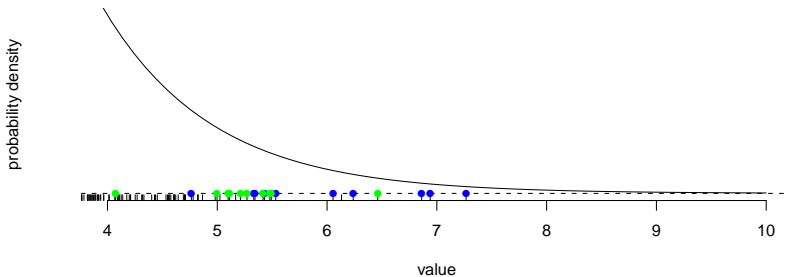
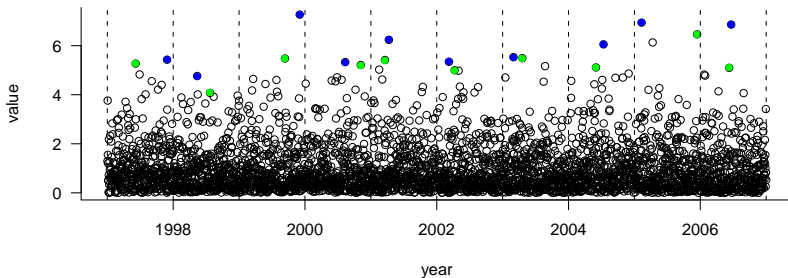




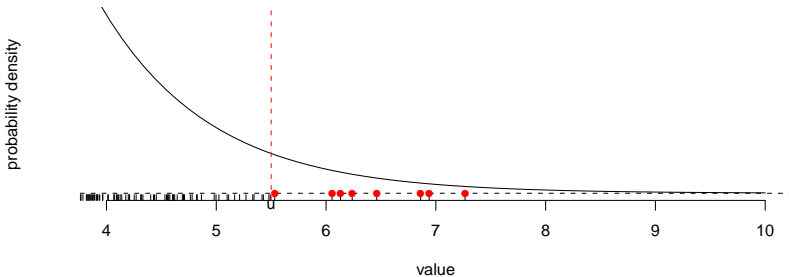
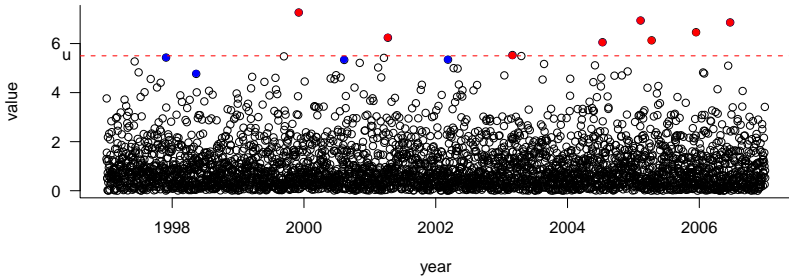
Block (annual) maxima



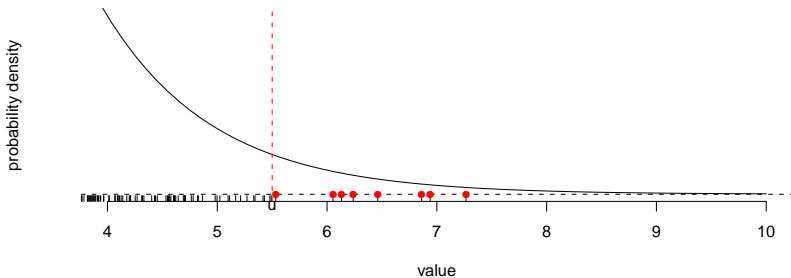
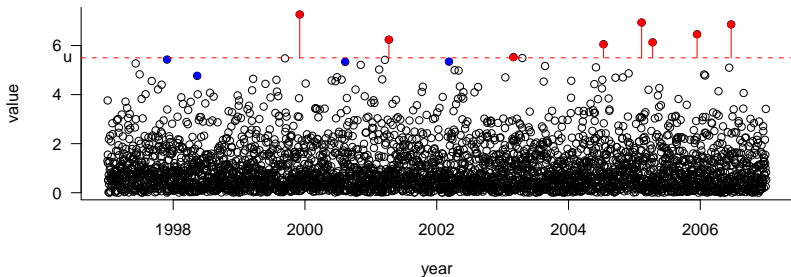
Maxima and second largest



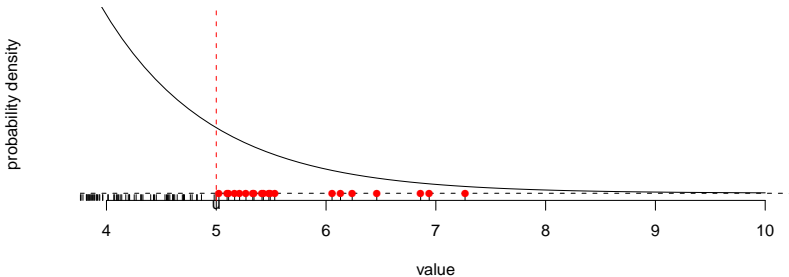
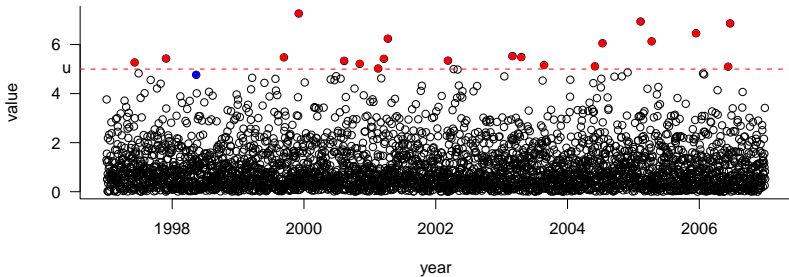
Threshold exceedances, $u=5.5$



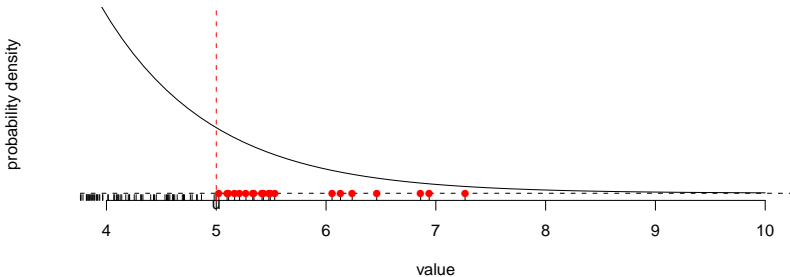
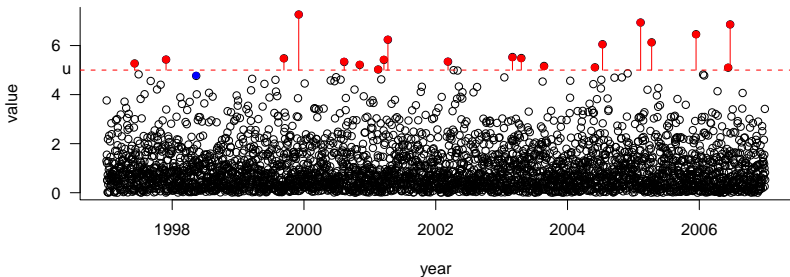
Exceedance amounts, $u=5.5$



Threshold exceedances, $u=5$



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- Any possible (non-degenerate) distribution of M_n as $n \rightarrow \infty$ is in the **GEV (Generalised Extreme Value)** family, with c.d.f.

$$G_{GEV}(z; \mu, \sigma, \xi) = \exp \left\{ - \left[1 + \xi \left(\frac{z - \mu}{\sigma} \right) \right]^{-1/\xi} \right\},$$

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- Suggests GEV distn as a model for, say, **annual maxima**.
- Upper end point is finite for $\xi < 0$ and infinite for $\xi \geq 0$.
- Related asymptotic model for r -largest order statistics.

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If Y_1, Y_2, \dots are independent then we get a **binomial-GP** model.

The GEV parameters (μ, σ, ξ) and the binomial-GP parameters (ρ_u, σ_u, ξ) are related ...

- $\sigma_u = \sigma + \xi(u - \mu)$;
- $\rho_u \approx \frac{1}{\lambda} \left[1 + \xi \left(\frac{u - \mu}{\sigma} \right) \right]^{-1/\xi}$.

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- Reparameterise: $(\rho_u, \sigma_u, \xi) \rightarrow (\mu, \sigma, \xi)$.

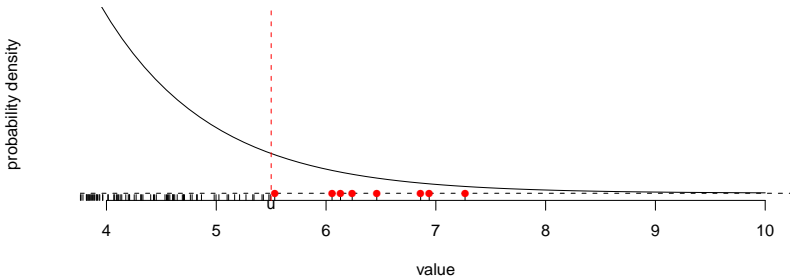
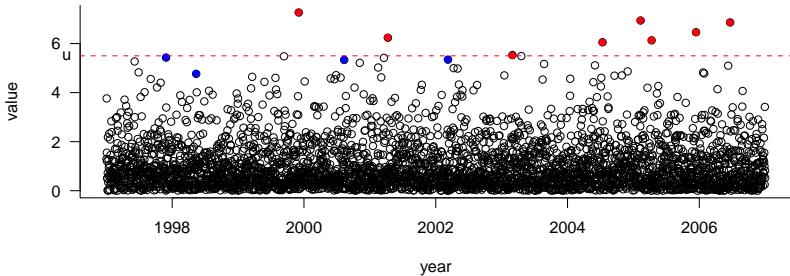
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- Poisson process \approx binomial process.

Threshold exceedances, $u=5.5$



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- Appeal to standard theory conditional on the covariates.
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- The PP model has the advantage (over the GP model) that its parameters are invariant to u .

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Possible approaches

1. **Model** the temporal dependence.
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Not necessary for the waves data because separate cluster (storm) are provided.

... but we will use 3. to adjust for **spatial dependence** .

- (Conditional) return levels.

Suppose that annual maxima are $\text{GEV}(\mu(x), \sigma(x), \xi(x))$

The *m*-year return level y_m satisfies

$$G_{\text{GEV}}(y_m; \mu(x), \sigma(x), \xi(x)) = 1 - \frac{1}{m},$$

so that

$$y_m = \mu(x) + \sigma(x) \left\{ \left[-\log \left(1 - \frac{1}{m} \right) \right]^{-\xi(x)} - 1 \right\}.$$

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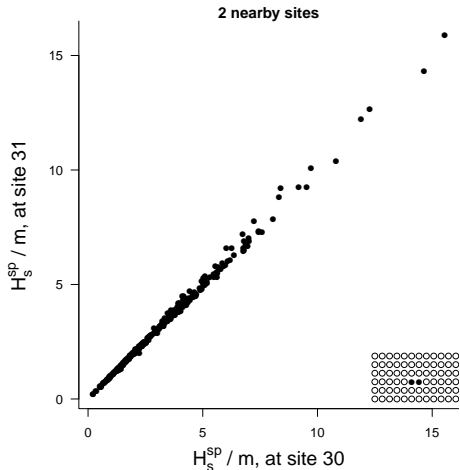
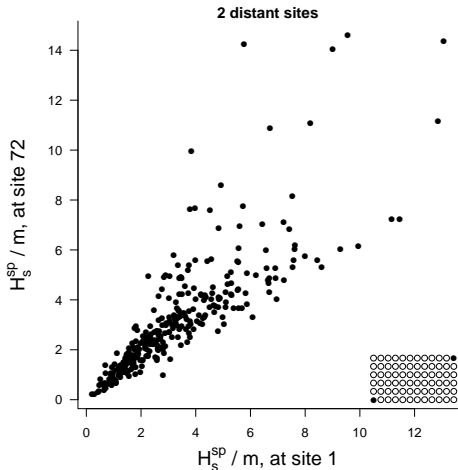
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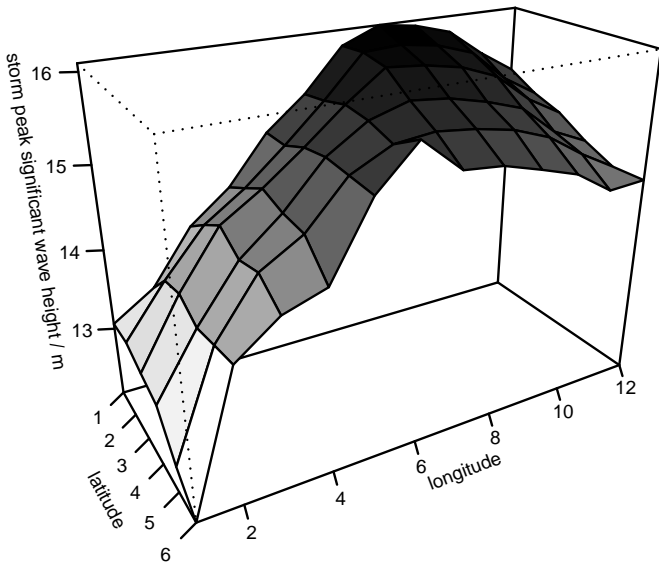
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- More specific quantities, e.g. probability that a certain value is exceeded during the next m years.





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- Spatial dependence not modelled explicitly: effect on marine structure is at one location.
- Storm peak from separate storms approximately independent.
- Time trends not apparent.
- Estimate return levels at centre of the region.

- **Spatial non-stationarity** : model spatial effects on EV parameters as Legendre polynomials in longitude and latitude.

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- **Threshold** : use **quantile regression** to achieve approx. constant probability p of threshold exceedance over space:
Model $100(1 - p)\%$ quantile as a function of covariates.
- **Spatial dependence** .
 - Estimate parameters assuming conditional independence of responses given covariate values.
 - Adjust standard errors etc. for spatial dependence.

Conditional on covariates \mathbf{x}_{ij} exceedances over a high threshold $u(\mathbf{x}_{ij})$ follow a **2-dimensional non-homogeneous Poisson process**.

If responses $Y_{ij}, i = 1, \dots, 72$ (space), $j = 1, \dots, 315$ (storm) are conditionally independent:

$$L(\theta) = \prod_{j=1}^{315} \prod_{i=1}^{72} \exp \left\{ -\frac{1}{\lambda} \left[1 + \xi(\mathbf{x}_{ij}) \left(\frac{u(\mathbf{x}_{ij}) - \mu(\mathbf{x}_{ij})}{\sigma(\mathbf{x}_{ij})} \right) \right]_+^{-1/\xi(\mathbf{x}_{ij})} \right\} \\ \times \prod_{j=1}^{315} \prod_{i: y_{ij} > u(\mathbf{x}_{ij})} \frac{1}{\sigma(\mathbf{x}_{ij})} \left[1 + \xi(\mathbf{x}_{ij}) \left(\frac{y_{ij} - \mu(\mathbf{x}_{ij})}{\sigma(\mathbf{x}_{ij})} \right) \right]_+^{-1/\xi(\mathbf{x}_{ij}) - 1} .$$

λ : mean number of observations per year;

$\mu(\mathbf{x}_{ij}), \sigma(\mathbf{x}_{ij}), \xi(\mathbf{x}_{ij})$: GEV parameters of annual maxima at \mathbf{x}_{ij} ;

θ : vector of all model parameters:

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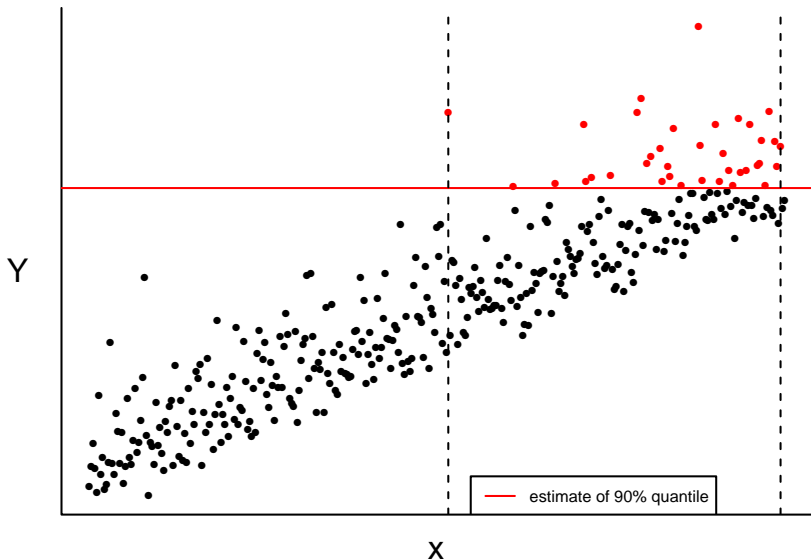
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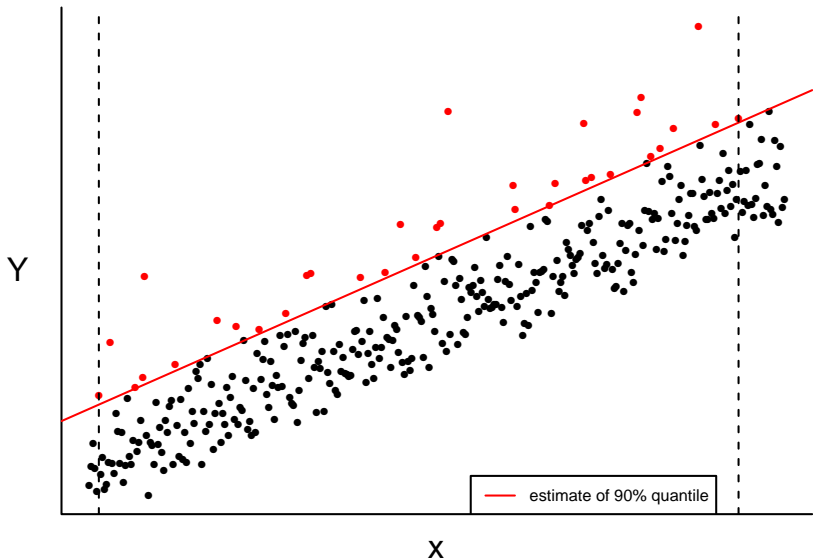
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Set $u(\mathbf{x}_{ij})$ so that $P(Y > u(\mathbf{x}_{ij}))$, is approx. constant for all \mathbf{x}_{ij} .

- Set $u(\mathbf{x}_{ij})$ by trial-and-error or by discretising \mathbf{x}_{ij} , e.g. different threshold for different locations, months etc.
- **Quantile regression (QR)** : model quantiles of Y as a function of covariates.





Let $p(\mathbf{x}_{ij}) = P(Y_{ij} > u(\mathbf{x}_{ij}))$. Then, if $\xi(\mathbf{x}_{ij}) = \xi$ is constant,

$$p(\mathbf{x}_{ij}) \approx \frac{1}{\lambda} \left[1 + \xi \left(\frac{u(\mathbf{x}_{ij}) - \mu(\mathbf{x}_{ij})}{\sigma(\mathbf{x}_{ij})} \right) \right]^{-1/\xi}.$$

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The form of $u(\mathbf{x}_{ij})$ is determined by the extreme value model:

- if $\mu(\mathbf{x}_{ij})$ and/or $\sigma(\mathbf{x}_{ij})$ are linear in \mathbf{x}_{ij} : **linear QR** ;
- if $\log(\mu(\mathbf{x}_{ij}))$ and/or $\log(\sigma(\mathbf{x}_{ij}))$ is linear in \mathbf{x}_{ij} : **non-linear QR** .

Data-generating process: for covariate values x_1, \dots, x_n

$$Y_i | X = x_i \stackrel{\text{indep}}{\sim} \text{GEV}(\mu_0 + \mu_1 x_i, \sigma, \xi).$$

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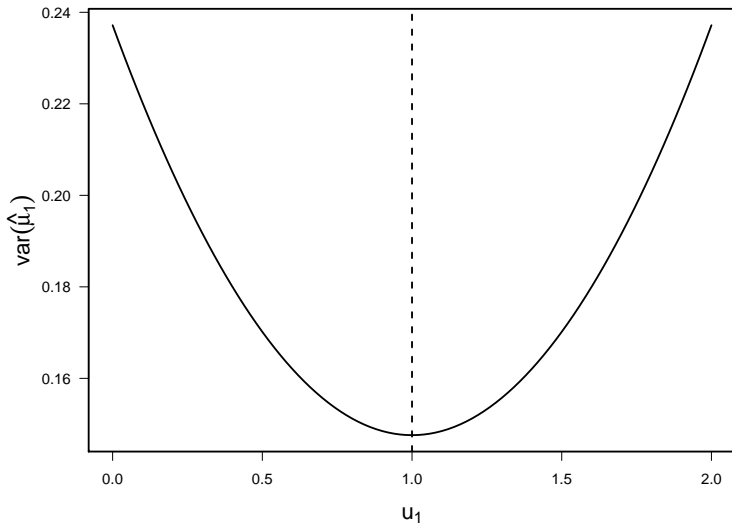
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- Calculate Fisher expected information for $(\mu_0, \mu_1, \sigma, \xi)$.
- Invert to find asymptotic V-C of MLEs $\hat{\mu}_0, \hat{\mu}_1, \hat{\sigma}, \hat{\xi}$ and hence $\text{var}(\hat{\mu}_1)$.
- Find the value of u_1 that minimises $\text{var}(\hat{\mu}_1)$.

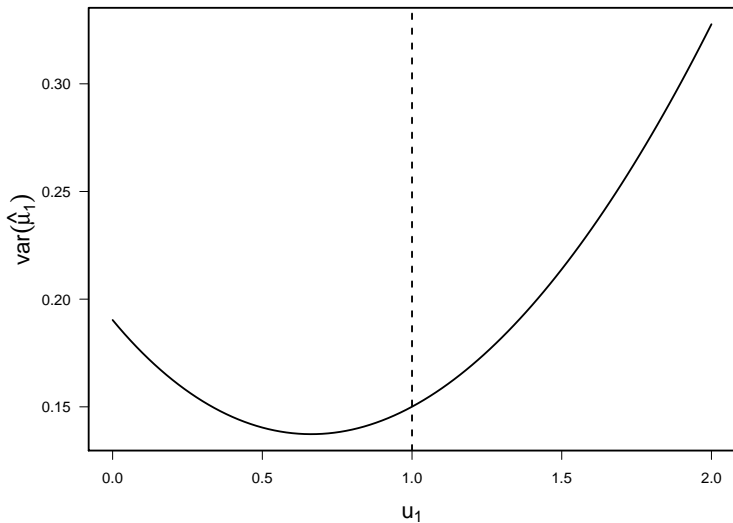
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Extensions:

- More general models.
- Effect of model mis-specification due to low threshold.

Independence log-likelihood:

$$l_{IND}(\theta) = \sum_{j=1}^k \sum_{i=1}^{72} \log f_{ij}(y_{ij}; \theta) = \sum_{j=1}^k l_j(\theta).$$

(storms) (space)

In regular problems, as $k \rightarrow \infty$,

$$\hat{\theta} \rightarrow N(\theta_0, H^{-1} V H^{-1}),$$

- $H =$ expected Hessian: $E \left(\frac{\partial^2}{\partial \theta^2} l_{IND}(\theta_0) \right)$;
- $V = \text{var} \left(\frac{\partial}{\partial \theta} l_{IND}(\theta) \right)$

- \hat{H} = observed Hessian, at $\hat{\theta}$;
- $\hat{V} = \sum_{j=1}^k U_j(\hat{\theta})^T U_j(\hat{\theta}), \quad U_j(\theta) = \frac{\partial l_j(\theta)}{\partial \theta}.$

Let $\hat{H}_A = \left(-\hat{H}^{-1} \hat{V} \hat{H}^{-1}\right)^{-1}.$

Chandler and Bate (2007):

$$I_{ADJ}(\theta) = I_{IND}(\hat{\theta}) + \frac{(\theta - \hat{\theta})' \hat{H}_A (\theta - \hat{\theta})}{(\theta - \hat{\theta})' \hat{H} (\theta - \hat{\theta})} \left(I_{IND}(\theta) - I_{IND}(\hat{\theta}) \right),$$

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- Adjust $I_{IND}(\theta)$ so that its Hessian is H_A at $\hat{\theta}$ rather than H .
- Preserves the usual asymptotic distribution of the likelihood ratio statistic.

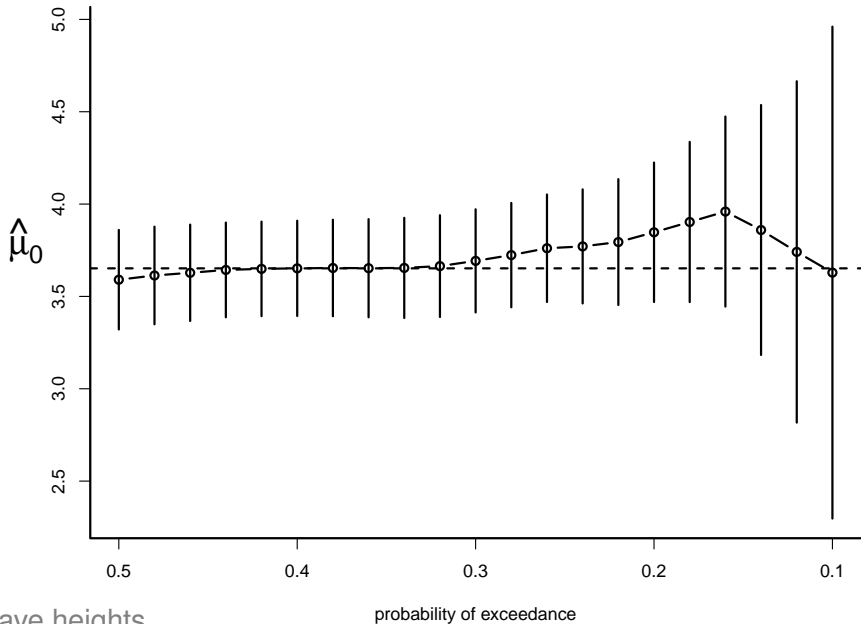
- Data simulated with spatial and temporal dependence and with spatial variation.
- Slight underestimation of standard errors : uncertainty in threshold ignored.

- Data simulated with spatial and temporal dependence and with spatial variation.
- Slight underestimation of standard errors : uncertainty in threshold ignored.
- Uncertainties in covariate effects of threshold are negligible compared to the uncertainty in the level of the threshold.
- Estimates of regression effects from QR and EV models are very close : both estimate extreme quantiles from the same data.
- To a large extent fitting the EV model accounts for uncertainty in the covariate effects at the level of the threshold.

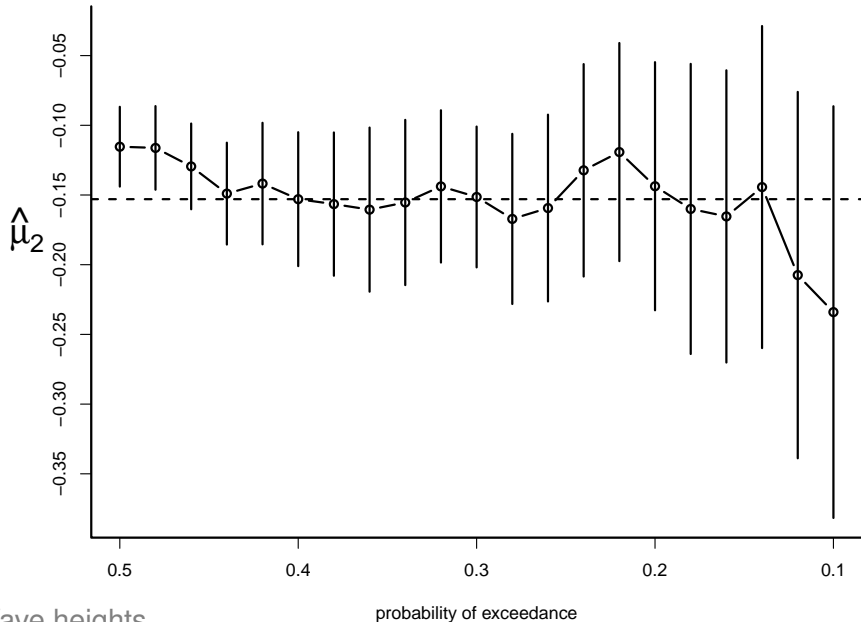
Threshold selection:

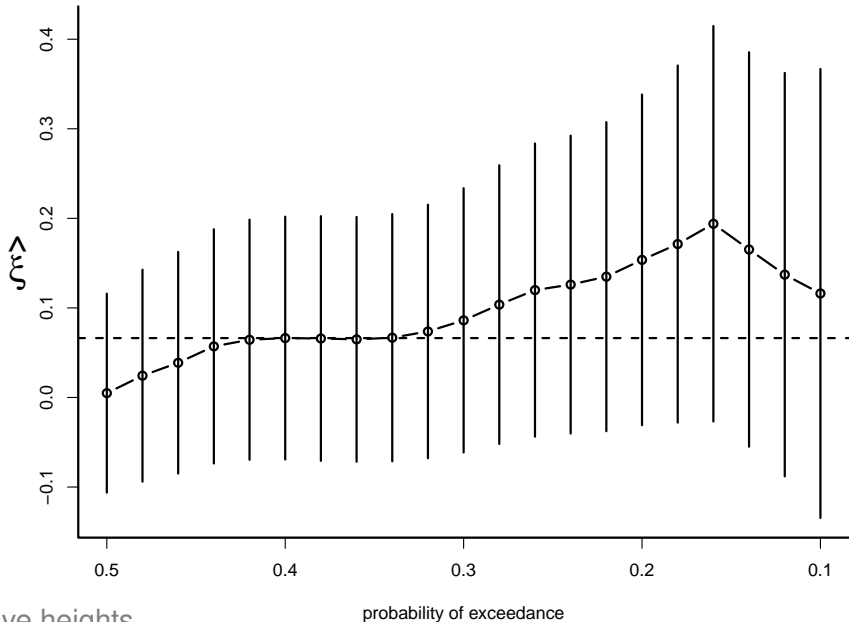
- Iterative: form of threshold depends on model.
- For given EV model set threshold using appropriate QR model.
- Choice of exceedance probability p : look for stability in parameter estimates.
- Based on μ (and u) quadratic in longitude and latitude, σ and ξ constant . . .

Threshold selection : μ intercept

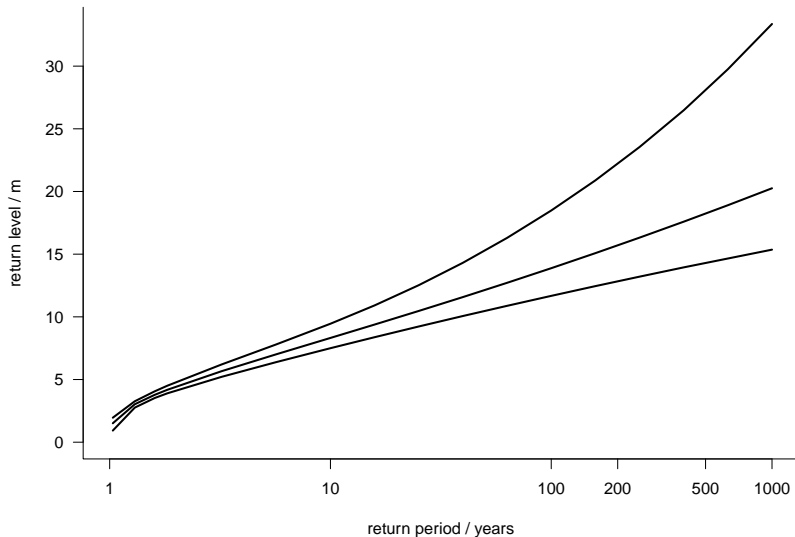


Threshold selection : μ coeff of latitude

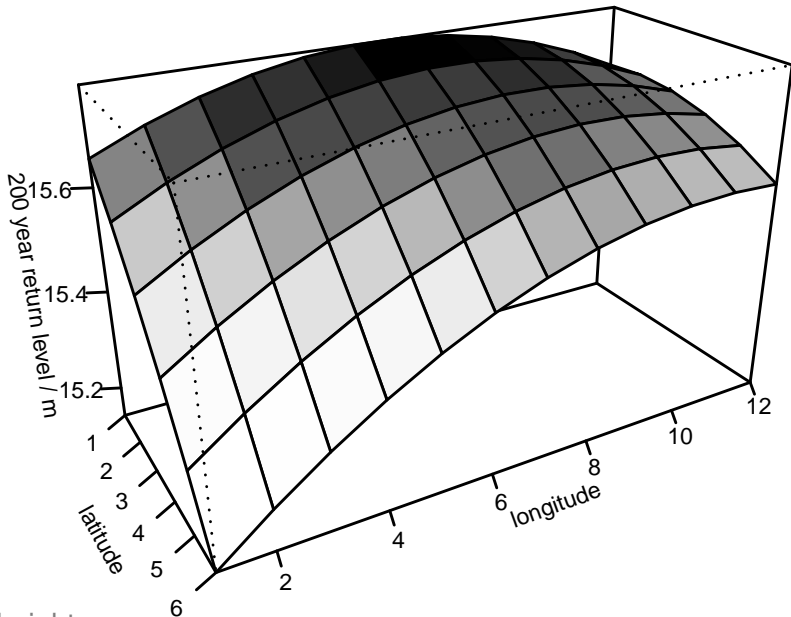




- Choice of p : look for stability in parameter estimates.
Use $p = 0.4$.
- Model diagnostics : slight underestimation at very high levels, but consistent with estimated sampling variability.
- QR model and EV model agree closely.
- $\hat{\xi} = 0.066$, with 95% confidence interval $(-0.052, 0.223)$.
- Estimated 200 year return level at centre of region is 15.70m with 95% confidence interval $(12.82, 22.20)$ m.



At-site 200 year return levels



Quantile regression

- a simple and effective strategy to set thresholds for non-stationary EV models;
- supported by simulation study;
- theoretical work is on-going;

Kysely *et al.* (2010): QR for time-dependent thresholds (GP).

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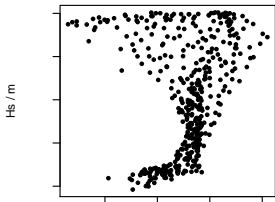
Kysely *et al.* (2010): QR for time-dependent thresholds (GP).

Comments

- Simpler r -largest order statistic analysis preferable for these data - but not for irregularly-spaced covariates.
- Other covariates: water depth; wave direction; 'hurricane alley'?
- $(\hat{U}\hat{H}^{-1})^T\hat{U}\hat{H}^{-1}$ more comp. stable than $\hat{H}^{-1}\hat{V}\hat{H}^{-1}$.

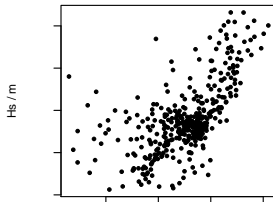
At-site order stats of H_s vs. water depth

largest



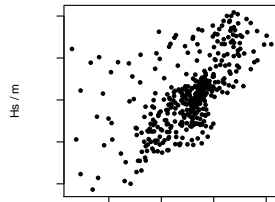
depth / m

2nd largest



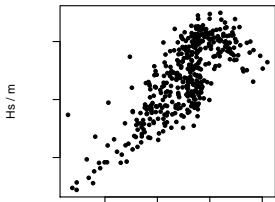
depth / m

3rd largest



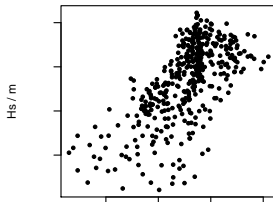
depth / m

4th largest



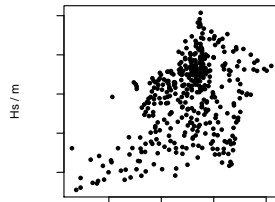
depth / m

5th largest



depth / m

6th largest



depth / m

Chandler, R. E. and Bate, S. B. (2007) Inference for clustered data using the independence loglikelihood. *Biometrika* **94** (1), 167–183.

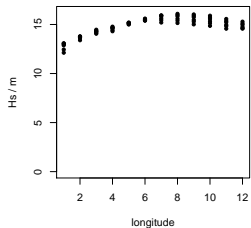
Fawcett, L. and Walshaw, D. (2007). Improved estimation for temporally clustered extremes. *Environmetrics*, **18**, 173188.

Kysely, J., Picek, J. and Beranová, R. (2010) Estimating extremes in climate change simulations using the peaks-over-threshold method with a non-stationary threshold *Global and Planetary Change*, **72**, 55-68.

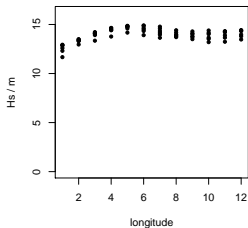
Northrop, P. J. and Jonathan, P. Threshold modelling of spatially-dependent non-stationary extremes with application to hurricane-induced wave heights. Published online in *Environmetrics*. (With discussion.)

Thank you for your attention.

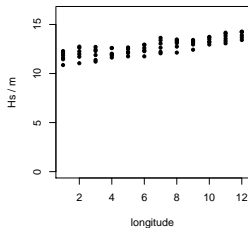
largest



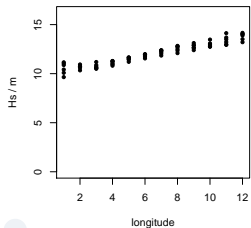
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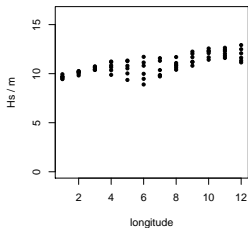
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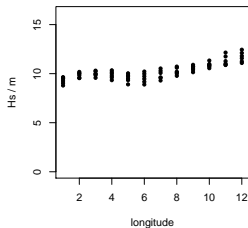
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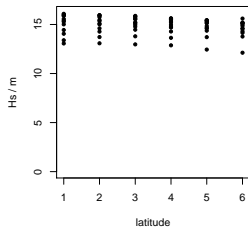
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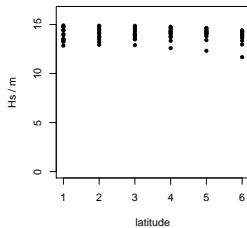
6th largest



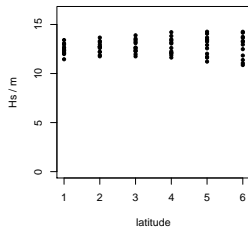
largest



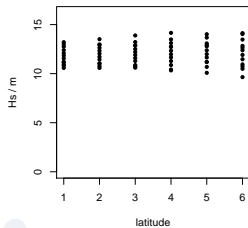
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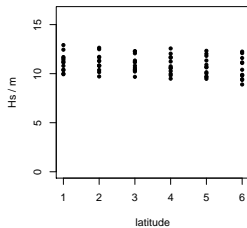
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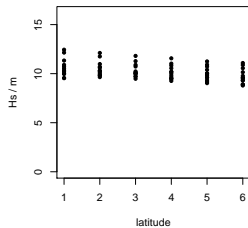
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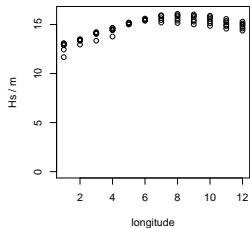
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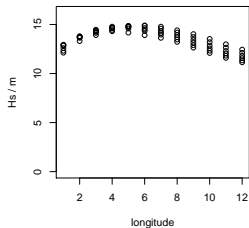
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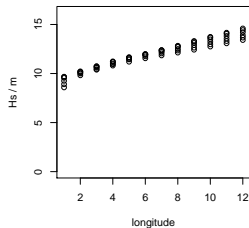
storm 314



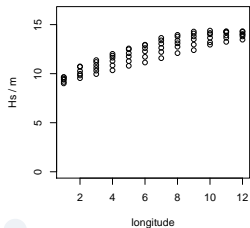
storm 45



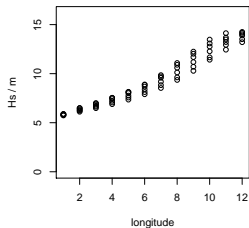
storm 311



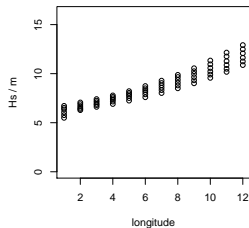
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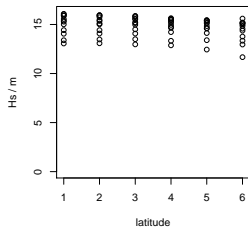
storm 220



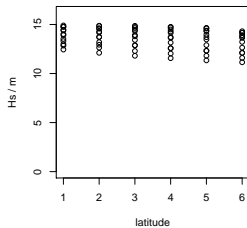
storm 278



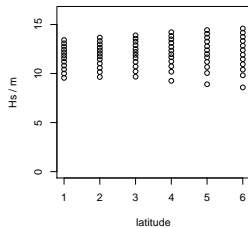
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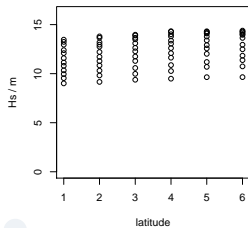
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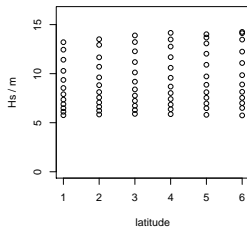
storm 311



storm 210



storm 220



storm 278

